



# Snow Modeling Capabilities within GSSHA





# Why Do We Care About Snow?

- Water Supply in the Western United States
  - Meltwater from the snowpack in the headwaters can provide 50-80% of the annual downstream water supply (Wahl, 1992).
  - Headwater catchments compose less than 25% of the total land area, but snowmelt from these areas provide roughly 70% of the annual runoff (Barros & Lettenmaier, 1993).
- Flooding







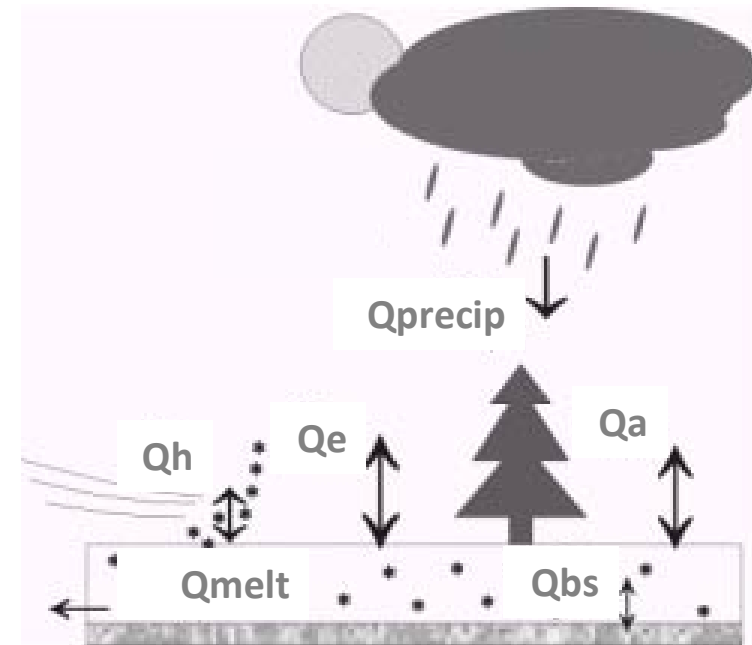
- Modeling of Snow Water Equivalent (SWE)
  - Four Snow Accumulation / Melt Methods
    - Energy Balance Method
    - Temperature-Index Method
    - Hybrid Energy Balance Method
    - Radiation-derived Temperature-Index Method
  - Accounting for the Snow Pack Dynamics within the Snow Pack
- Adjustments to HMET Forcing Data
- Melt Water Transport
  - Vertical Flow through Snow Pack
  - Lateral Flow through Snow Pack
  - Frozen Ground to Impede Infiltration
- Examples





# Four Snow Methods in GSSHA

- Energy Balance Method (EB) – accounts for the energy fluxes between the snow pack, ground layer, and atmosphere. If enough energy is input into the pack then melting occurs.
- Temperature-Index Method (TI) – The amount of melt within the snow pack is based on the temperature, precipitation, and calibrated parameters. Based on SNOW-17.
- Hybrid Energy Balance Method (HY, Default) – modification of the Energy Balance model that accounts for snow pack temperature dynamics (heat deficits).
- Radiation-Derived Temperature-Index Method (RTI) – Also based on SNOW-17, but uses a radiation-derived proxy temperature instead of air temperature in the melt equations.



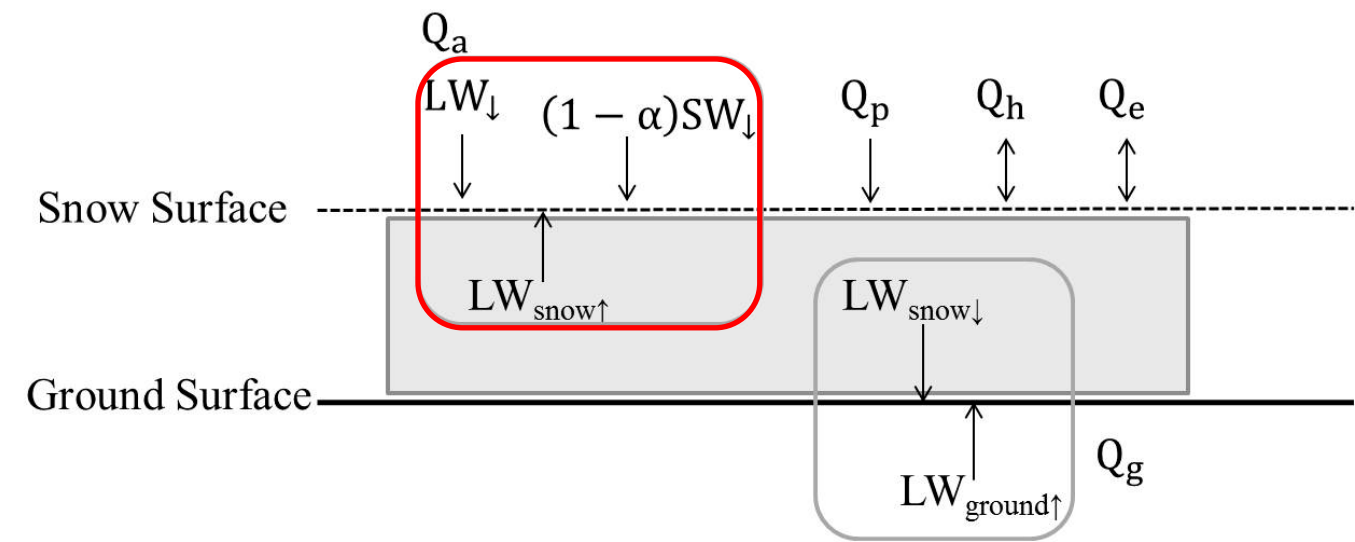
Picture Adapted from: Tarboton 1996



# Simulating SWE – EB vs TI (In General)

~~Temp Energy Bal~~ (EB)

$$M_{\text{melt}} = \frac{\sum Q}{L_f} dt$$



- $Q_a$  – Net Radiation
- $Q_p$  – Precipitation Heat Flux
- $Q_h$  – Sensible Heat Flux
- $Q_e$  – Latent Heat Flux
- $Q_g$  – Ground Heat Flux
- $\sigma$  – Stefan Boltzmann Constant
- $\varepsilon$  – Emissivity (0-1)
- $\alpha$  – Albedo
- $L_f$  – Latent Heat of Fusion



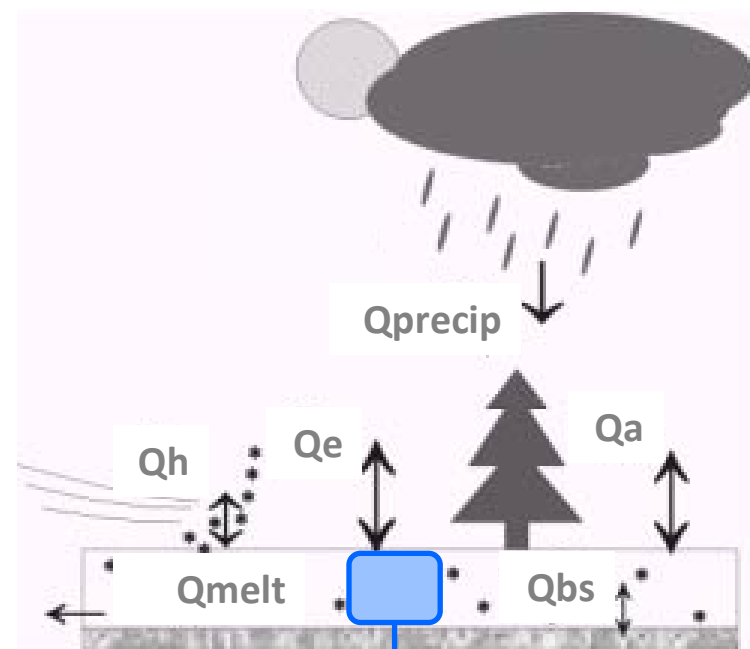


# Simulating SWE – EB & HY Methods

$$Q_{\text{melt}} = Q_a - Q_{\text{bs}} + Q_e + Q_h + Q_{\text{precip}}$$

- $Q_{\text{melt}}$  total energy available to melt snow
- $Q_e$  energy due to evaporation and sublimation  
f(temperature, humidity, wind)
- $Q_h$  sensible heat transfer due to turbulence  
f(temperature, pressure, wind)
- $Q_a$  energy due to longwave radiation  
f(temperature)
- $Q_{\text{bs}}$  energy due to longwave emission by soil  
considered constant ( $27 \text{ cal cm}^{-2} \text{ hr}^{-1}$ )
- $Q_{\text{precip}}$  energy due to precipitation  
f(precipitation, temperature)

Assumes 1 cc of snow will melt for every 336 Joules  
which = 80 cal per gram of water



Picture Adapted from: Tarboton 1996

HY Method accounts for cold  
content (heat deficit) within  
the snowpack.





Based on SNOW-17 (Anderson 1968; 1973; and 2006)

- Melt occurs once heat deficit is overcome
- Melt rates ( $M$ ) based on Air Temperature and calibrated Melt Factors

$$M = [M_f \cdot (T_a - MBASE) + 0.0125 \cdot P \cdot f_r \cdot T_r] \quad \text{Melt under normal circumstances}$$

$$M = \sigma \cdot dt \cdot [(T_a + 273)^4 - 273^4] + 0.0125 \cdot P \cdot f_r \cdot T_r + 8.5 \cdot f_u \cdot (dt/6) \cdot [(rh \cdot e_{sat} - 6.11) + 0.00057 \cdot P_a \cdot T_a] \quad \text{Melt during rain events.}$$

where:

$M_f$  = melt factor, varies daily between a calibrated minimum and maximum value

$P$  = precipitation

$f_r$  = fraction of precipitation in form of rain

$f_u$  = wind function

$MBASE$  = temperature at which snow begins to melt







# Simulating SWE – TI to RTI Method

Temperature Index (TI)

$$Melt = M_f T_a dt$$

$$T_a \approx \left( \frac{LW_{\downarrow}}{\sigma \varepsilon_{air}} \right)^{1/4}$$

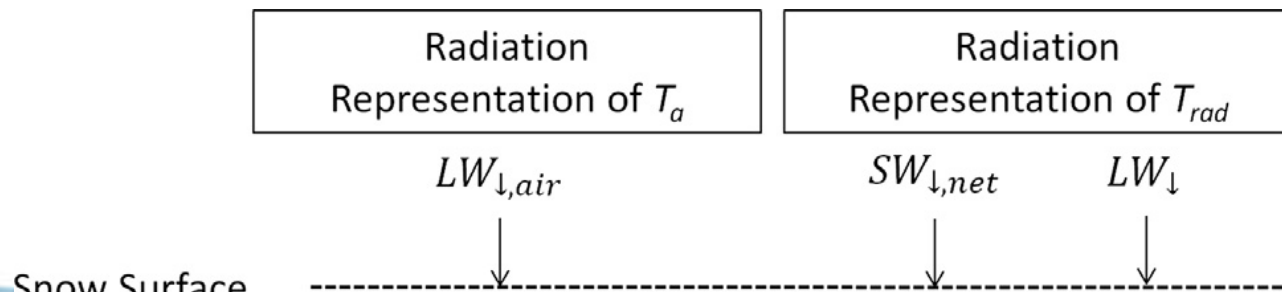
Requires:  
Temperature  
Precipitation

Radiation-Derived Temperature Index (RTI)

$$Melt = M_f T_{rad} dt$$

$$T_{rad} \approx \left( \frac{LW_{\downarrow} + SW_{\downarrow net}}{\sigma \varepsilon_{snow}} \right)^{1/4}$$

Requires:  
Temperature  
Precipitation  
Cloud Cover







# Simulating SWE – RTI Method

## How to get $T_{rad}$

- Account for SW radiation and related reductions

$$SW_{\downarrow} = S_0 K_r K_v K_c K_{atm} K_s K_t$$

$S_0$ =Solar Constant

$K_r$ =ratio of actual earth-sun distance to mean earth-sun distance (fraction)

$K_v$ =vegetation

$K_c$ =clouds

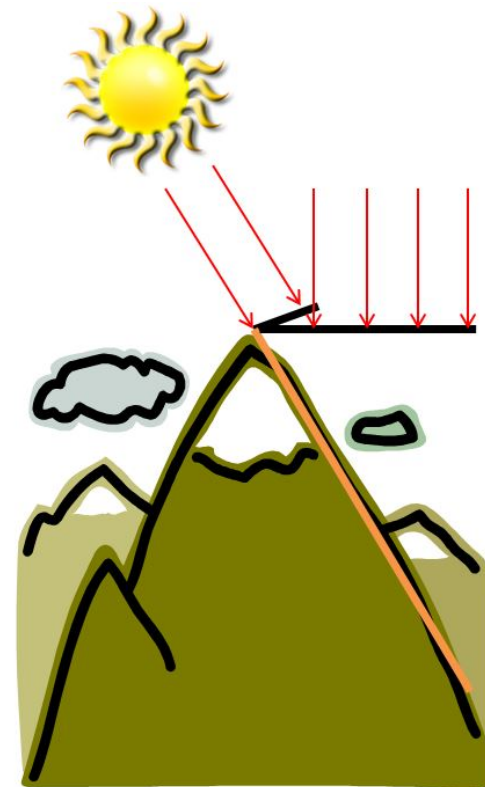
$K_{atm}$ =atmospheric absorption and scattering

$K_t$ =topographic shading

$K_s$ =aspect angle

- Include snow albedo model (Henneman & Stefan, 1999)
- Account for LW radiation from clouds and canopy ( $LW_{\downarrow}$ )
  - LW clear sky, including emissivity of air (Bras, 1990)
  - Increase of LW due to clouds (TVA, 1972)
  - Adjustment of LW due to vegetation (Liston & Elder, 2006).

For more details see Follum et al. (2015)





# Simulating SWE - Overview

- Energy Balance Method
  - Uses Energy Balance Algorithms to determine melt and accumulation, but does not include Heat Deficit.
  - Typically underestimates SWE when topography not accounted for in the model.
- Temperature-Index Method
  - Uses temperature, precipitation, and calibrated parameters to simulate snow accumulation and melt.
  - Incorporates the Heat Deficit / Snow Pack Dynamics
  - Requires Calibration (9 Parameters)
- Hybrid Energy Balance (Default in GSSHA)
  - Incorporates the Energy Balance Melting Algorithms
  - Incorporates the Heat Deficit / Snow Pack Dynamics (4 Parameters)
- Radiation-Derived Temperature-Index Method
  - Same as TI Method, but accounts for spatial heterogeneity in energy, and therefore produces a more accurate spatial representation of the snowpack.
  - Requires Calibration (8+ Parameters)





# Adjustments to HMET Forcing Data

- Three methods for adjusting HMET with Elevation
  - **SIMPLE**: Define a constant temperature lapse rate, then GSSHA adjusts temperature. Based on change in temperature, relative humidity and pressure can also be adjusted.
  - **COMPLEX**: Let GSSHA automatically adjust temperature, pressure, and relative humidity based on elevation using MALR.
  - **PUNT**: Input raster-based forcing data from an atmospheric model, such as MicroMet (Liston & Elder, 2006)

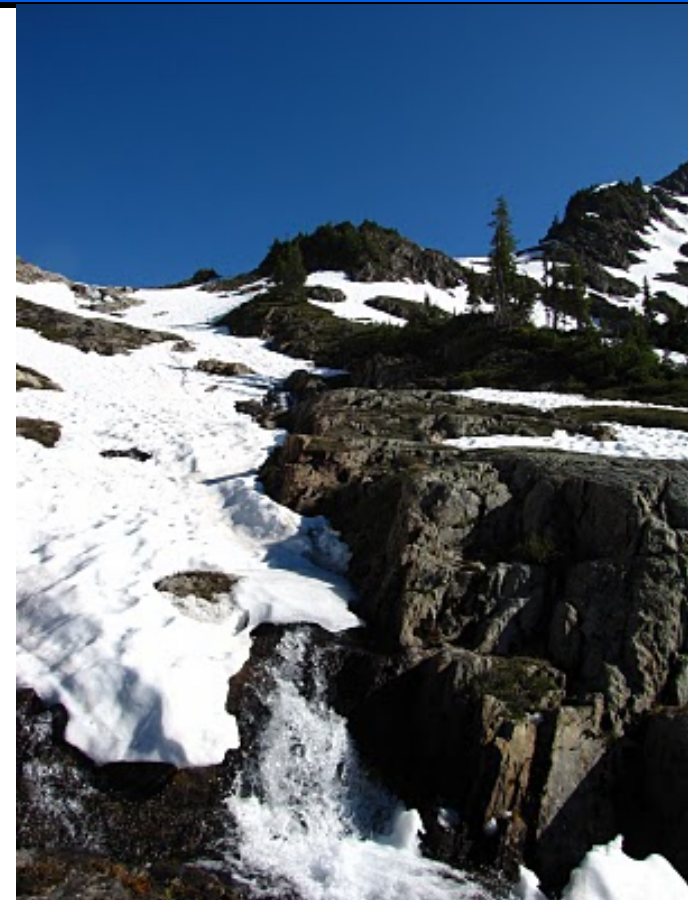
Typically, I use a simple Lapse Rate





# Melt Water Transport

- Once snow melts the water is transported through the system in several ways:
  - Vertically through the Snowpack ❄️
  - Laterally through the Snowpack ❄️
  - Infiltration into Groundwater ❄️
  - Overland Flow
  - Channel Flow, Etc.
- GSSHA simulates the vertical and lateral flow through the Snowpack
- Accurate groundwater simulations is also very important for capturing the timing of flows. A frozen ground simulation determines when the ground is frozen - inhibiting infiltration into groundwater.

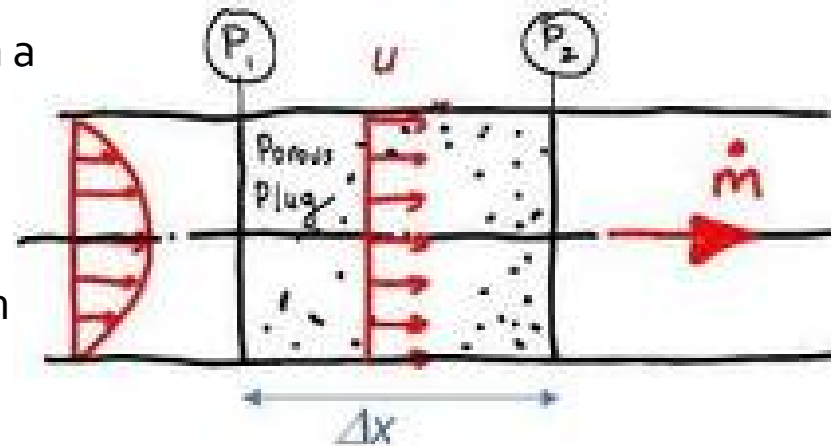






# Melt Water Transport – Basics

- Flow through Snow is the same as a flow through a porous medium.
- A form of Darcy's Equation is typically used to determine flux rates through the snow pack, both vertically and laterally.
- Vertical Flow is typically considered Unsaturated Flow.
- Lateral Flow at the bottom of the snow pack is typically considered Saturated Flow.
- GSSHA uses the SNAP model (Albert 1998) to determine the saturation, saturated / unsaturated hydraulic conductivity, and effective porosity of the snow pack in each cell.



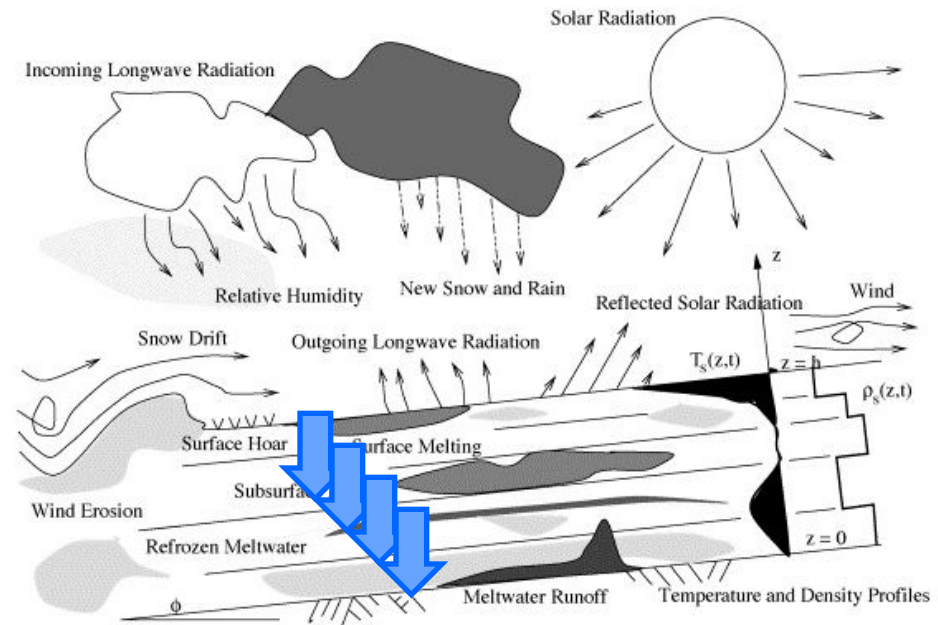


# Melt Water Transport – Vertical Flow

- Vertical Flow is considered unsaturated flow, but the hydraulic conductivity and effective porosity change with the degree of Saturation.

$$k_w = kS^n$$

- Flow is simulated through a single snow layer, but can have multiple wave fronts through the pack – based on both Albert (1998) and Bengtsson (1981).



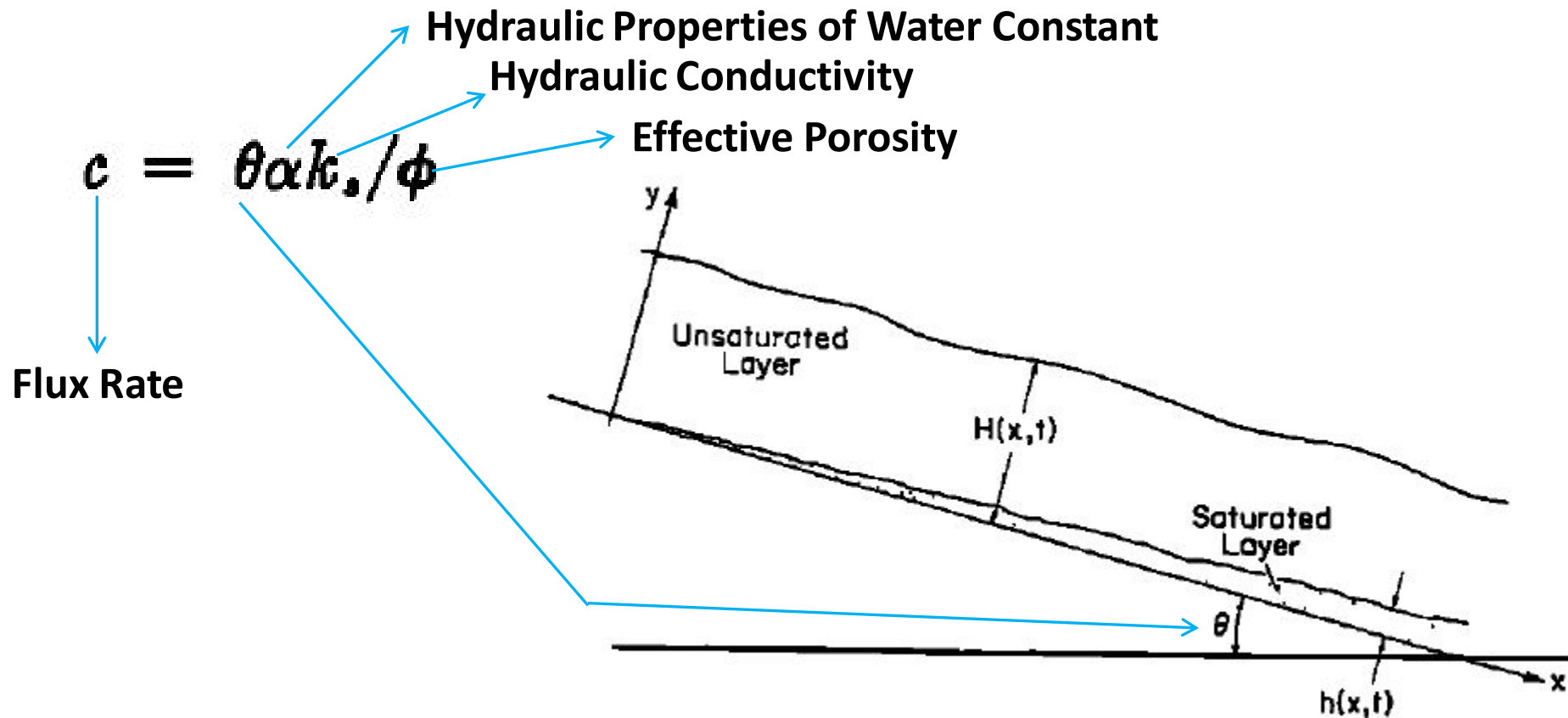
From: Bartelt (2002)





# Melt Water Transport – Lateral Flow

- Once the melt water reaches the ground it is considered Saturated Darcian Flow and uses methods developed by Colbeck (1974) to determine the flux volumes between each cell.



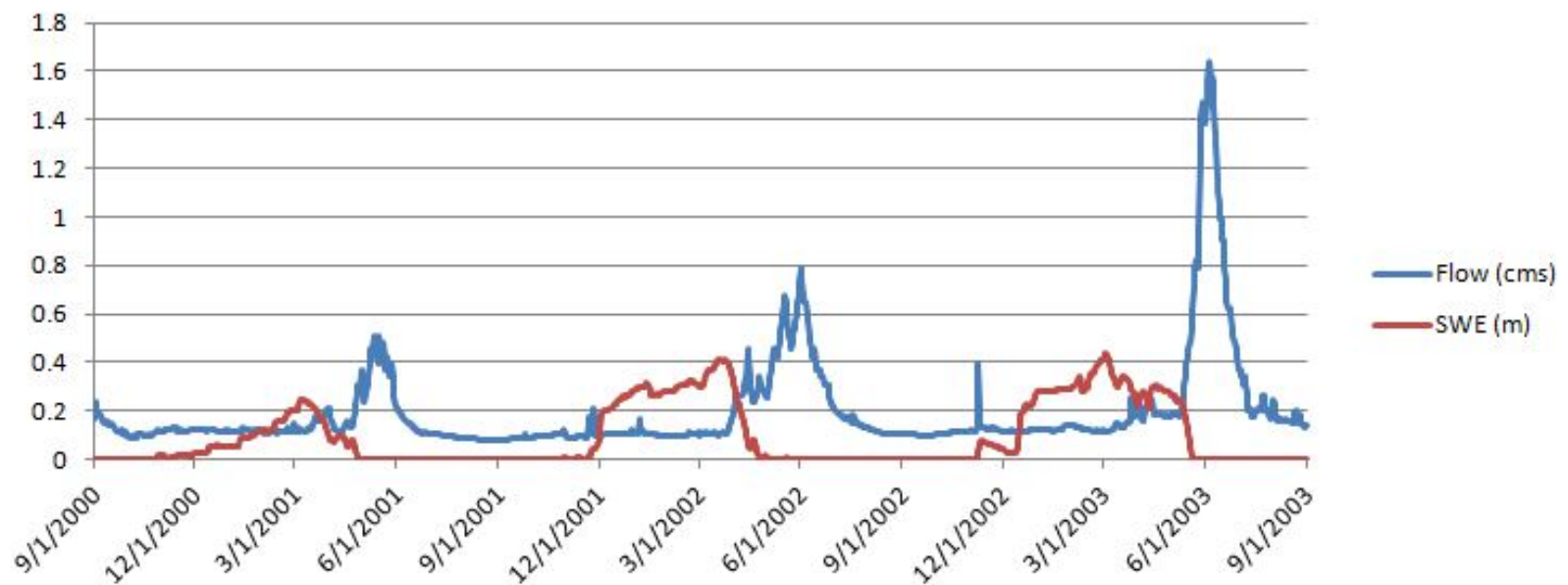
From: Colbeck (1974)



# Melt Water Transport – Frozen Ground

- GSSHA uses a temperature-index method (CFGI model; Molnau & Bissel, 1983) to determine when the ground beneath a snow pack is frozen, thus preventing infiltration.
  - See for more info: [http://www.gsshawiki.com/Frozen\\_Soil:Frozen\\_Soil](http://www.gsshawiki.com/Frozen_Soil:Frozen_Soil)
- The Handbook of Snow (Gray and Male, 1981) states that for long-term sustained water yields the groundwater flow component may be most important aspect considered.

## Trout Creek Basin, CA







# Melt Water Transport - Overview

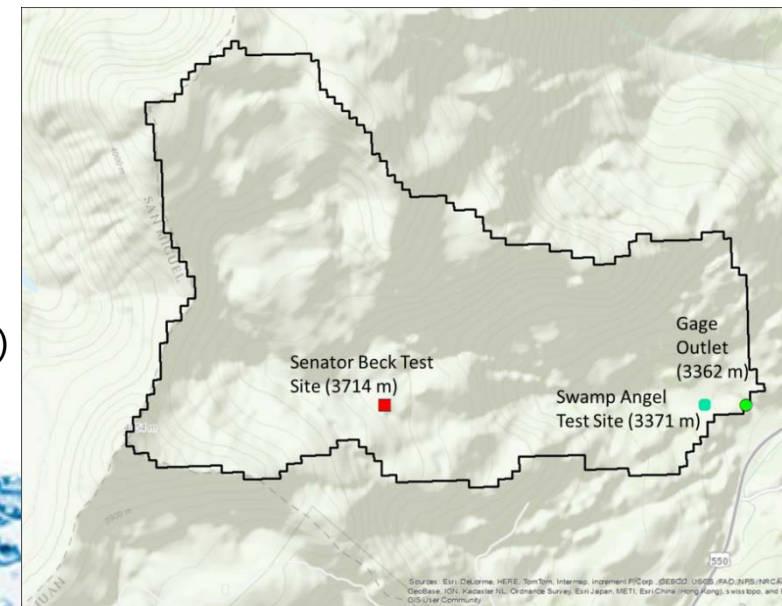
- Vertical Flow through Snow Pack
- Lateral Flow through Snow Pack
- Frozen Ground Simulation
- Existing routing mechanisms already in GSSHA





# Test Basin: Senator Beck Basin, CO (SBB)

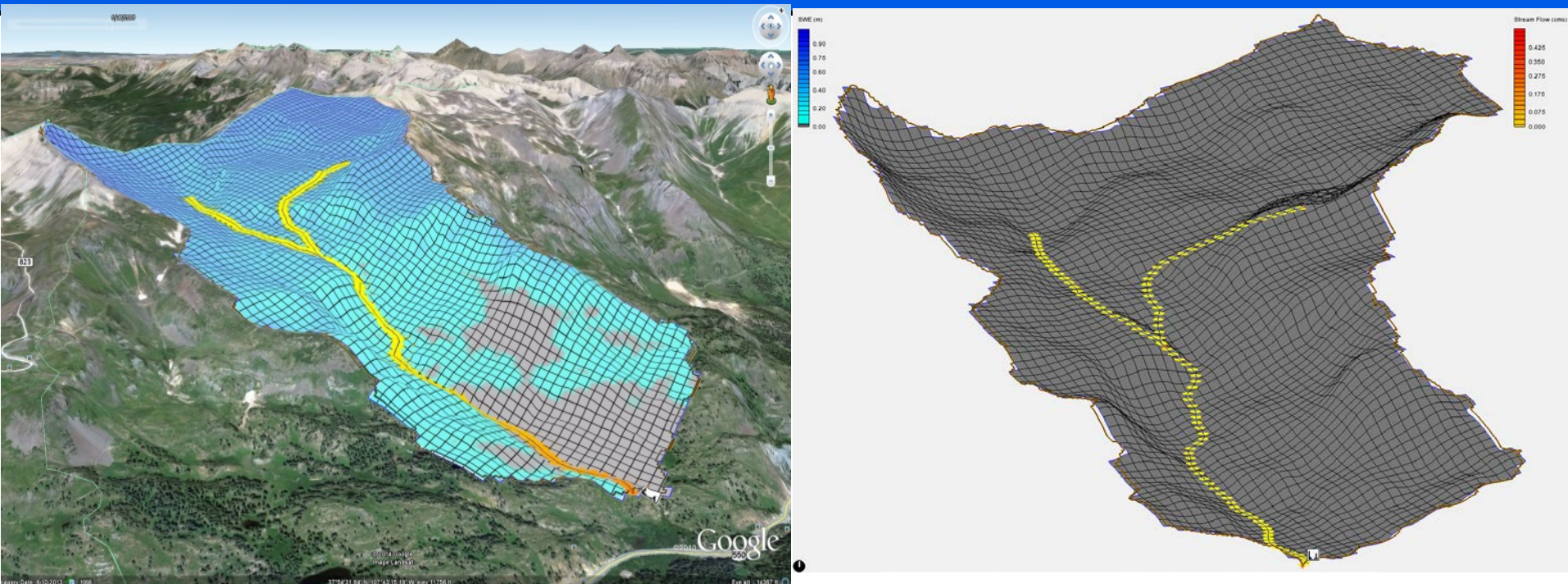
- Test Basin: Senator Beck Basin, CO
  - Drainage Area: 2.91 km<sup>2</sup>, Elev: 3362 – 4118 m
  - Alpine terrain with primarily bare rock and tundra, with some forest below 3600-m
- Point Data Sources
  - Hydrometeorology at 3 sites from WY 2003 through current (Landry et al., 2014)
    - Temperature, precipitation, relative humidity, pressure, and wind speed
    - SWE, soil moisture, and shortwave (SW) radiation
  - Cloud cover data from Telluride Regional Airport (~16 km to northwest)
  - Streamflow data at outlet (Landry et al., 2014)
- Spatial Data Sources
  - SCA estimates from LandSat Imagery (31 images)
    - Spectral signatures processed using ERDAS Imagine
    - Assigned classifications: snow, no snow, and snow fringe
  - Elevation -> 1/3 Arc Second NED (Gesch et al., 2002)
    - projected to 30 m grid
  - Land Cover -> 2006 NLCD (Fry et al., 2011)
  - Soils Data -> SSURGO dataset (Soil Survey Staff, 2014)







# Test Basin: Senator Beck Basin, CO (SBB)



## Highlights:

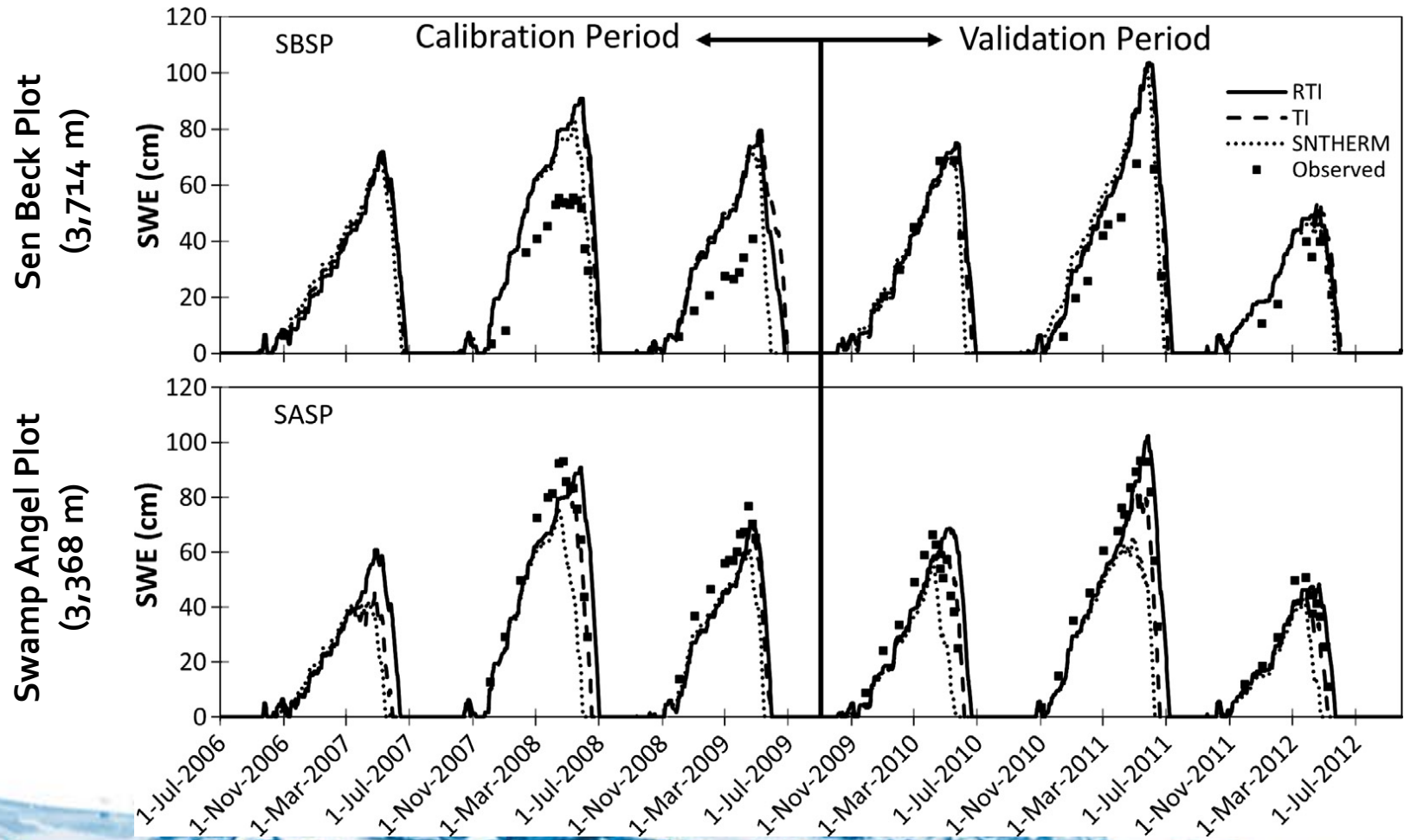
- Showed accurate snow simulation in an Alpine terrain (Follum et al., 2015)
- RTI model more accurate than TI and EB models





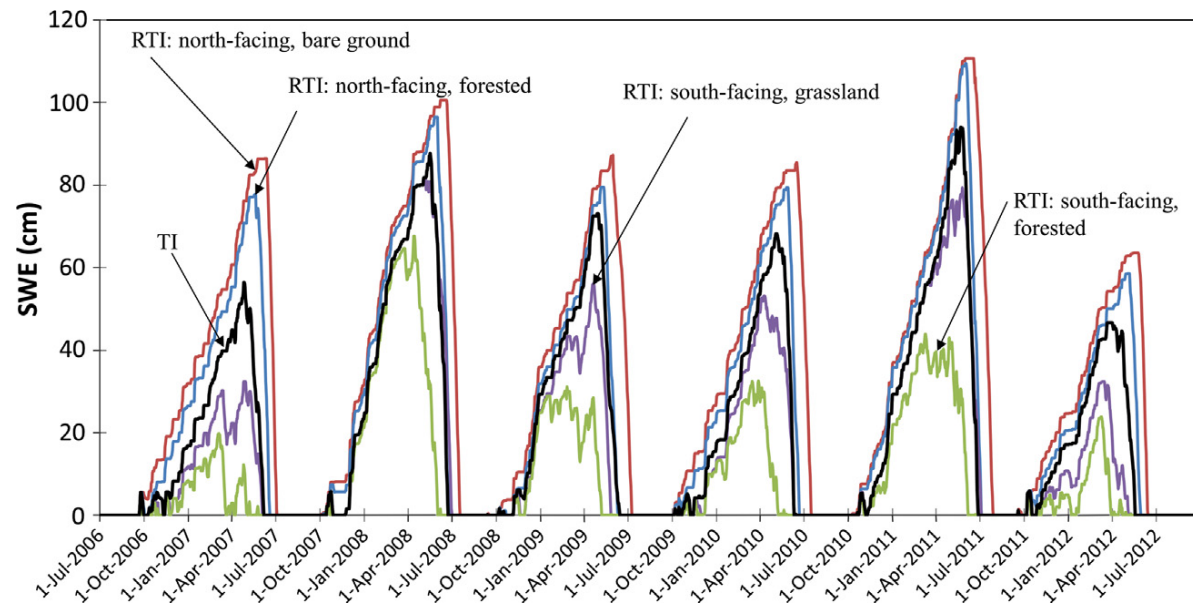
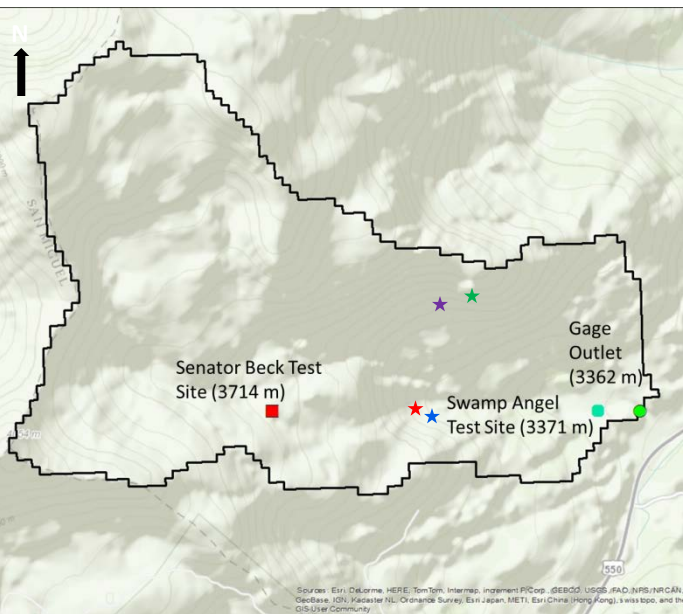
# SBB using TI and RTI Model

- Modest improvement in SWE at two gage sites



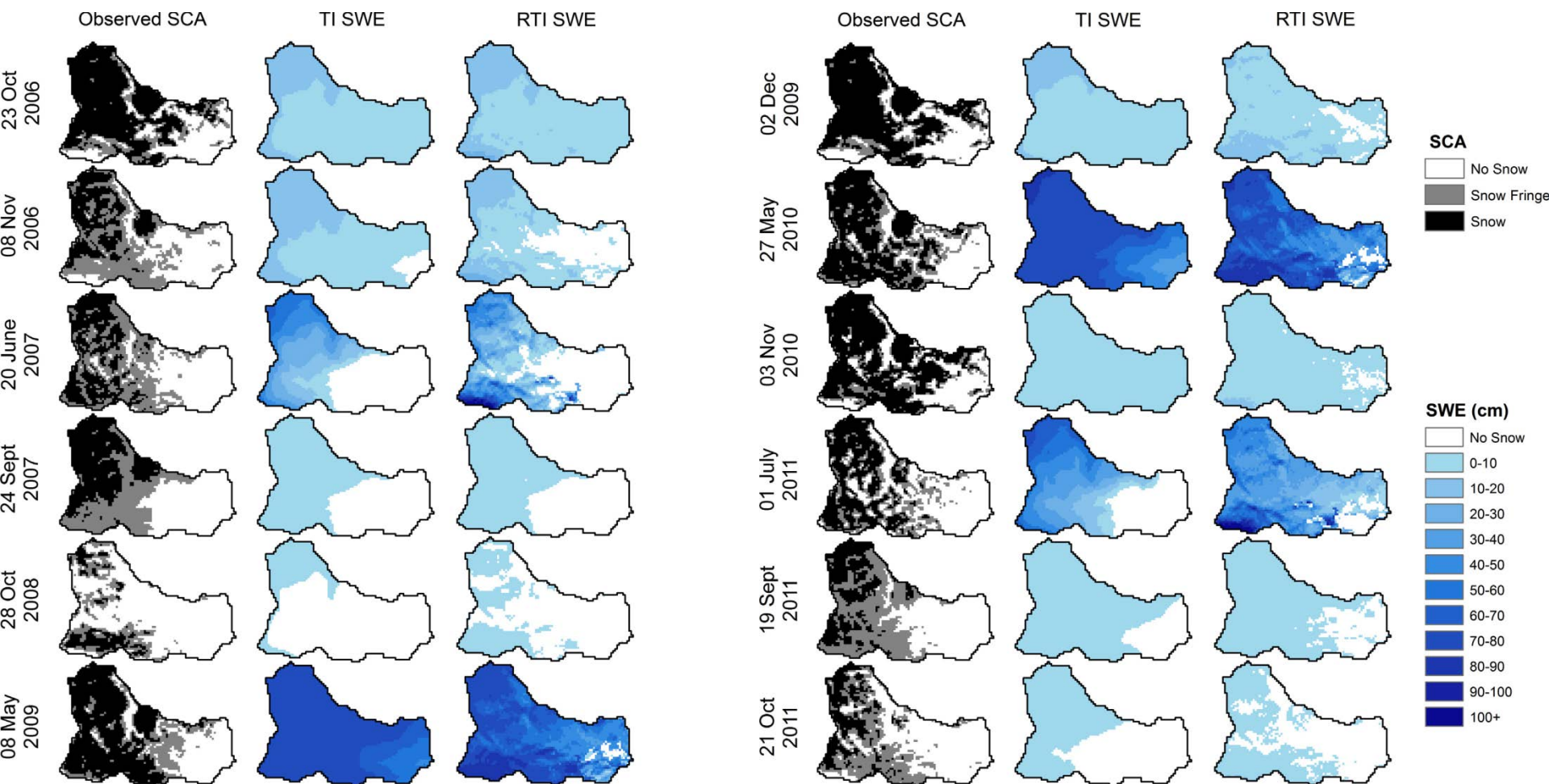


- **Variability in SWE due to topography and vegetation**



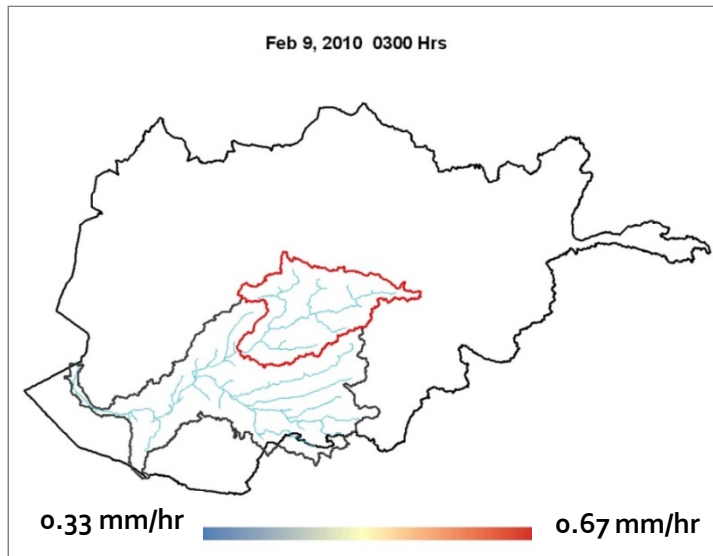


# SBB using TI and RTI Model



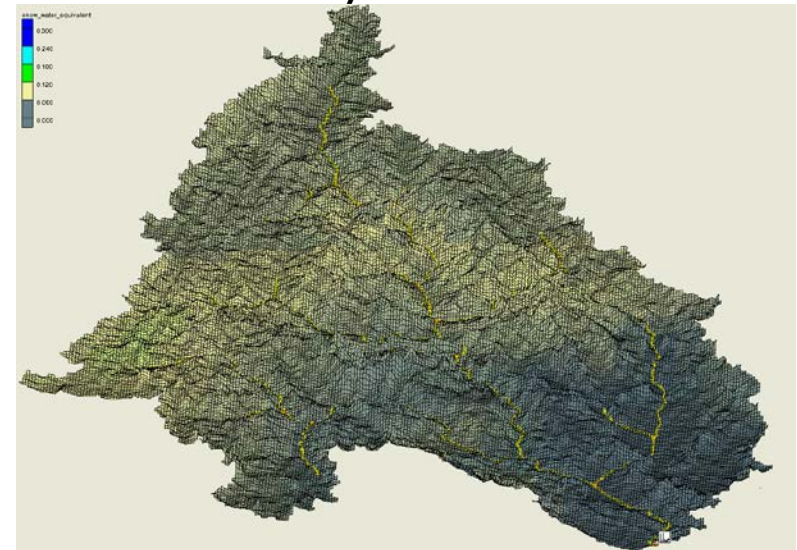


# Helmand River Basin, Afghanistan



- Watershed-based
- Manual setup
- Limited forecasting window

Upper Helmand Basin Simulation:  
January – June 2010







# Data Sources

- NRCS / USDA - <http://www.wcc.nrcs.usda.gov/>
  - SNOTEL, SCAN, Snowcoarses
- National Snow & Ice Data Center - <http://nsidc.org/>
- Remotely Sensed
  - Landsat, MODIS, AVHRR
- CZO's - <http://criticalzone.org/national/>
- Test Watersheds:
  - Senator Beck Basin, CO
  - Niwot Ridge, CO
  - Fraser Experimental Forest, CO
  - Loch-Vail, CO
  - Reynolds Experimental Watershed, ID
  - Sleepers River, VT
  - Hubbard Brook, NH
  - HJ Andrews, OR
  - Marmot Creek, Canada Rockies







# Questions?

## References:

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